

Research Article

Volume 19 Issue 1 - April 2019  
DOI: 10.19080/IJESNR.2019.19.556003

Int J Environ Sci Nat Res

Copyright © All rights are reserved by Antensay Mekoya

# Estimation of Evaporation using Daily and Ten-Minute Class-A Pan Data from Automatic Measuring Pressure Sensor Instrument at Tharandt, Germany



Antensay Mekoya<sup>1\*</sup>, Christian Bernhofer<sup>2</sup> and Uta Moderow<sup>2</sup>

<sup>1</sup>National Meteorology Agency (NMA), Ethiopia

<sup>2</sup>Technische Universität Dresden, Germany

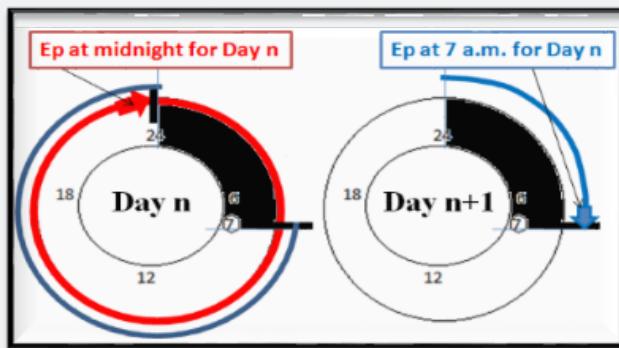
Submission: April 08, 2019; Published: April 30, 2019

\*Corresponding author: Antensay Mekoya, NMA of Ethiopia, Ethiopia

## Abstract

Evaporation is major component of global water cycle and water balance of a small/large irrigation area, reservoir or lake, and a catchment. In this study, evaporation from Class-A pan ( $E_p$ ) for the years from 2004 to 2013 at Tharandt, Germany is calculated and analyzed using daily and ten-minute Class-A pan water level data from automatic measuring pressure sensor instrument.

Daily  $E_p$  is first calculated at two different times; at 7 a.m. and at midnight. Because daily  $E_p$  calculated at 7 a.m. had shown less errors as it had fewer negative values ( $n = 43$  out of a total of 2145 values) than  $E_p$  calculated at midnight ( $n = 84$  out of a total of 1971 values);  $E_p$  calculated at 7 a.m. is selected for the calculation of  $E_p$ . The correlation between  $E_p$  calculated at midnight and  $E_p$  calculated at 7 a.m. was 'very good' ( $R^2 = 0.87$ ,  $MSE = 0.32 \text{ mm d}^{-1}$ ). Therefore,  $E_p$  calculated at midnight is used for filling as well as for correcting  $E_p$  calculated at 7 a.m. values. Accordingly, missing values of daily  $E_p$  at 7 a.m. are filled using  $0.908 \times \text{daily } E_p \text{ at midnight values} + 0.385$  (Eq. 7). In Eq. 7, the cause for non-zero offset (0.385) could be instrument error. Assuming no instrument error, out of 1836 days,  $E_{pd}$  which is the difference between 7 a.m. and midnight  $E_p$  (see Figure 1) was larger than zero for 1184 days (64.5% of the days). This means out of 1836 days, for 64.5% of the days next day 'night time'  $E_p$  was greater than its previous day 'night time'  $E_p$ . Also, for at least 54 days  $|E_{pd}| \leq 1.5 \text{ mm d}^{-1}$  which means that the 'night time' daily  $E_p$  had to be  $\geq 1.0 \text{ mm d}^{-1}$ .



**Figure 1:** Understanding  $E_{pd}$  which is daily  $E_p$  at 7 a.m. ('blue') minus daily  $E_p$  at midnight ('red') for day n.

Finally, 2098 daily values of  $E_p$  are calculated from March to November for the available data; however, only the summer half year (April to September,  $n=1702$ ) values of  $E_p$  are mainly used for most of the analysis.

Generally, the accuracy of the self-recording ten-minute and daily water level measurements from Class-A pan at Tharandt from 2004 to 2013 can be considered as very good. However, the measurement should be carefully checked as it might have sensitivity to other than pressure or water depth difference in the pan. Hence, the sensitivity of the pressure sensor instrument at 7 a.m. and at midnight for same pressure (depth of pan water) might have slight difference.

**Keywords:** Tharandt; Class-A pan Evaporation ( $E_p$ );  $E_p$  at 7 a.m. and at midnight,  $E_{pd}$

**Abbreviations:**  $E_p$ : Class-A Pan Evaporation ( $E_p$ );  $E_{pd}$ :  $E_p$  at 7 a.m. minus  $E_p$  at midnight; ET: Evapotranspiration; V: Water Level at 7 a.m.; P: Precipitation at 7 a.m.; V': Water Level at Midnight; RF: Precipitation at Midnight; NSE: Nash-Sutcliffe Efficiency; MAE: Mean Absolute Error; RMSE: Root Mean Square Error; RSR: RMSE-Observations Standard Deviation Ratio; MPE: Mean Percent of Error; 'night time': The Time between Midnight and 7 a.m.

## Introduction

Measured and estimated evaporation data has been in use by agricultural, hydrological, hydro meteorological, irrigation and soil and water conservation applications. Evaporation or evapotranspiration (ET) which is a major component of the global water cycle and the hydrologic budget or water balance of small or large irrigation area, reservoir or lake and a catchment is important consumer of energy. On average, across all continents

about 70% of precipitation reaching the land surface evaporates; in dry regions (e.g., Australia) this ratio is higher and can reach up to 90% and in Europe to approximately 60% of the annual rainfall (Nova'k 2012 & Baumgartner and Reichel, 1975) [1,2].

This article describes the methods used for estimation of evaporation from Class-A pan in a very humid temperate region (Tharandt, Germany).



**Figure 2:** Site view of Tharandt Meteorological Station (Photo taken on March 21, 2017 at 15:32)

## Materials and Methods

The station is situated in flat and grass covered area and fulfils the WMO standard for meteorological measurements. This station has special characteristics because it is situated at the bottom of a V-shaped valley and close to an asphalt road, buildings and the Weißeritz River (see Figure 2 & 3). Therefore, it has a reduced sky view factor and consequently sunshine may comparatively reaches the area late in the morning and leaves earlier in the evening resulting in lower sunshine hours. Due to high shelter effects at low level, it is expected that the wind speed at 2m (which was derived from the wind speeds at 3m and at 10m) would have been higher than real values (if actual measurement had been conducted at 2m height).

The pan used for measurement of pan evaporation is a standard Class-A pan evaporimeter (see Figure 3). The readings of the water level in the pan for every 10-minute interval ( $V'$ ) as well as for daily basis ( $V$ ) are recorded automatically by a pressure sensor device.

### Daily Class-A pan Evaporation at 7 a.m. and at midnight

Using Eq. 4 and using daily water level ( $V$ ) and the corresponding precipitation at 7 a.m. ( $P$ ), daily pan evaporation from Class-A pan ( $E_p$ ) is calculated at 7 a.m. Similarly,  $E_p$  at midnight is calculated using 10-minute water level ( $V'$ ) and daily precipitation at midnight (RF).

The change in water level for a day say d ( $\Delta V'_d$ ) at midnight is calculated by subtracting the water level at day d at 23:59:00 hour ( $V'_{d,1}$ ) from the water level at previous day (d-1) at 23:59:00 hour ( $V'_{d-1,1}$ ); i.e.,  $\Delta V'_d = V'_{d-1} - V'_{d,1}$ .  $\Delta V'_d$  can also be calculated by taking the sum of 144 ten-minute water level differences for each day. Both will give same result if  $V'$  has no missing values. Similarly, the change in water level for a day say d ( $\Delta V_d$ ) at 7 a.m. is calculated by subtracting the next day (day d+1) water level at 7 a.m. ( $V_{d+1,1}$ ) from the water level at day d at 7 a.m. ( $V'_{d,1}$ ); i.e.,  $\Delta V_d = V_d - V_{d+1,1}$ . Note that those days with missing  $V$  or  $V'$  data are excluded in the calculation of  $\Delta V$  and  $\Delta V'$ .

## Class-A pan Evaporation ( $E_a$ ) estimation

According to Dingman [3], pan evaporation is calculated using eq. 1 below:



**Figure 3:** Standard Class-A evaporation pan at Tharandt Met. Station (Photo source: IHM, TU Dresden).

Where,

$E$  pan evaporation (in mm)

$P$  precipitation (in mm) during  $\Delta t$ ,

$V_1$  &  $V_2$  the storages (in mm) at the beginning and end of  $\Delta t$ , respectively.

According to (WMO 1994 Sec. 9.2), the amount of evaporation that has occurred between two observations of water level in the pan ( $E$ ) is calculated using Eq. 2 below:

$$E = P \pm \Delta d \quad (\text{WMO 1994 Sec. 9.2}) \quad \dots \dots \dots \quad (2)$$

Where,

$P$  the depth of precipitation during the period between the two measurements,

$\Delta d$  the depth of water added (+) to or removed (-) from the pan.

Combining Eq. 1 and Eq. 2, we get:

$$E = P - (V_2 - V_1) \pm \Delta d \quad \dots \dots \dots \quad (3)$$

Modifying Eq. 3 and replacing  $E$  with  $E_{\text{pan}}$  or  $E_p$ , we have:

$$E_p = P + (V_1 - V_2) \pm \Delta d \quad \dots \dots \dots \quad (4)$$

Where,  $E_p$  is daily evaporation computed as the difference in Class-A pan water level on successive days, corrected for any precipitation and  $\Delta d$  during the period.  $P$ ,  $\Delta d$ ,  $V_1$  and  $V_2$  are as defined above. For the calculation of daily  $E_p$ , Eq. 4 is used throughout this article with  $\Delta d = 0$  (because of missing  $\Delta d$  values). However, 41 days have information for 'special features' ('Besonderheiten\_Daten') like emptying or drawing out (pumping off) some water, cleaning (e.g. using Anti-Algae

chemicals) and refilling of pan (including information about confirmation of no precipitation during the refilling time); removal (fishing out) of grass, seeds, coarse dirt or suspended matter (solids) from the pan; and so forth which are excluded from the calculation of  $E_p$ .

Note that the maximum possible value of daily  $E_{\text{pan}}$  in  $\text{mm d}^{-1}$  can be set to be equal to the upper estimate of daily PET in  $\text{mm d}^{-1}$  ( $\text{PET}_{\text{max}}$ ) (see Eq. 6).  $\text{PET}_{\text{max}}$  is estimated to be the extreme maximum value of the ratio of daily net radiation (in  $\text{MJ m}^{-2} \text{d}^{-1}$ ) calculated using Eq. 19 and (daily) latent heat of vaporization (in  $\text{MJ kg}^{-1}$ ) calculated using Eq. 5. It is calculated with the assumption that all the available energy provided by radiation is consumed (used for vaporization).

$$\lambda = 2.501 - (2.361 \cdot 10^{-3}) T \quad [4] \quad \dots \dots \dots \quad (5)$$

Where,

$\lambda$  latent heat of vaporization (in  $\text{MJ kg}^{-1}$ )

$T$  air temperature (in  $^{\circ}\text{C}$ )

$$E_{p_{\text{max}}} = \text{PET}_{\text{max}} = \text{MAX} = \left( \frac{R_n}{\lambda} \right) \quad \dots \dots \dots \quad (6)$$

Where,  $E_{p_{\text{max}}}$  is the maximum possible upper limit of  $E_p$  (in  $\text{mm d}^{-1}$ ),  $\text{PET}_{\text{max}}$  is the maximum possible upper limit of PET (in  $\text{mm d}^{-1}$ ),  $\lambda$  latent heat of vaporization (in  $\text{MJ kg}^{-1}$ ) and  $R_n$  net radiation (in  $\text{MJ m}^{-2} \text{d}^{-1}$ ).

Note, however, that the estimated net radiation used in this paper is based on measurements above grass level and that the Class-A pan will have a different radiation balance. Therefore, Eq. 6 can be a suited check for PET estimations according to Haude, Wendling, and Penman and  $\text{ET}_o$  but only a rough check for  $E_p$ .  $E_{p_{\text{max}}}$  resulted 7.198 ( 7.2)  $\text{mm d}^{-1}$ .

**Table 1:** Example for filling missed daily values of  $E_p$  (for 1<sup>st</sup> and 2<sup>nd</sup> of April) using ten-year daily averages of  $E_p$  for the available ten years.

Year	2004	2005	2006	2007	2008	2009	2010	2011	2012	2013	Average
Actual and Average of daily $E_p$ data for the available ten years (in $\text{mm d}^{-1}$ )											
1-Apr				2.32			1.7	1.94	1.55		1.88
2-Apr		2.19		2.27	1.72	1.73	0.74	2.24	0.77		1.67
Actual daily $E_p$ data (in $\text{mm d}^{-1}$ ); missing values filled using average values											
1-Apr	1.88	1.88	1.88	2.32	1.88	1.88	1.7	1.94	1.55	1.88	1.88
2-Apr	1.67	2.19	1.67	2.27	1.72	1.73	0.74	2.24	0.77	1.67	1.67

Monthly  $E_p$  is calculated by aggregation of daily  $E_p$  values for months from April to October of each year. Note however that few days at the beginning of April and at the end of October had considerable missing daily  $E_p$  values. Hence, the average values of daily  $E_p$  for the available ten years (from 2004 to 2013) is used for filling missing daily  $E_p$  values of each year (see Table 1).

## Results and Discussion

### Comparing Class-A pan evaporation at 7 a.m. and at midnight

Daily values of pan evaporation from Class-A pan ( $E_p$ ) were calculated at 7 a.m. and at midnight using Eq. 4. The calculation

was performed for  $n = 2145$  and 1971 days {from daily  $P$  and RF data and their corresponding fully available daily ( $V$ ) and ten-minutes ( $V'$ ) water level data, respectively}.

Generally daily  $E_p$  at 7 a.m. resulted in greater values than daily  $E_p$  at midnight. It also resulted in less (121) negative values whereas daily  $E_p$  at midnight resulted in more (213) negative values (see Table 2). Thus, comparatively  $E_p$  at 7 a.m. has the advantage of having a greater number of values with a smaller number of negative values which made it preferable to  $E_p$  at midnight. Therefore,  $E_p$  calculated at 7 a.m. is selected in this article. Its missing values are filled using  $E_p$  at midnight values following the description in section 3.2.

**Table 2:** Comparison of 7a.m. and midnight daily  $E_p$  values at Tharandt over ten years for three ranges of values (cases) and before application of any correction.

Case A: for all values of $E_p$	$E_p$ at 7 a.m.	$E_p$ at midnight	$E_{pd}^a$
Number of values (n)	2145	1971	1966 <sup>b</sup>
Number of days with $E_p \geq 7.2 \text{ mm d}^{-1}$	28	21	5
Number of days with $E_p \leq -0.5 \text{ mm d}^{-1}$	43	84	85
Number of days with $E_p < 0 \text{ mm d}^{-1}$	121	213	686
Extreme maximum value [ $\text{mm d}^{-1}$ ]	77.52	79.43	13.66
Extreme minimum value [ $\text{mm d}^{-1}$ ]	-77.88	-78.4	-41.37
Average [ $\text{mm d}^{-1}$ ]	2.32	2.16	0.21
Case B: for $-0.5 \text{ mm d}^{-1} \leq E_p \leq 7.2 \text{ mm d}^{-1}$	$E_p$ at 7 a.m.	$E_p$ at midnight	$E_{pd}^a$
Number of values (n)	1998 (1836) <sup>c</sup>	1866 (1836) <sup>c</sup>	1836
Average [ $\text{mm d}^{-1}$ ]	2.23 (2.23) <sup>d</sup>	2.03 (2.03) <sup>d</sup>	0.2
Case C: for $0 \leq E_p \leq 7.2 \text{ mm d}^{-1}$	$E_p$ at 7 a.m.	$E_p$ at midnight	$E_{pd}^a$
Number of values (n)	1871 (1699) <sup>c</sup>	1738 (1699) <sup>c</sup>	1699
Average [ $\text{mm d}^{-1}$ ]	2.37 (2.39) <sup>d</sup>	2.21 (2.21) <sup>d</sup>	0.19

<sup>a</sup> $E_{pd}$  is used to denote daily  $E_p$  at 7 a.m. minus daily  $E_p$  at midnight

<sup>b</sup>out of the 1966 days for two days both  $E_p$  at 7 a.m. and  $E_p$  at midnight had equal values

<sup>c</sup>the number in parenthesis is the number of days where both  $E_p$  at 7 a.m. and  $E_p$  at midnight had values;

<sup>d</sup>the corresponding average value.

To evaluate the error among case A, case B, and case C of Table 2, the methods for comparison and evaluation of models which are discussed and applied by considering the daily  $E_p$  at

midnight as observation values (i.e.,  $x$  or  $O_i$ ) and daily  $E_p$  at 7 a.m. as model (estimated) values (i.e.,  $y$  or  $P_i$ ).

**Table 3:** Brief statistical summary for comparison of daily  $E_p$  at 7 a.m. and at midnight for three ranges of values (cases) before application of any correction.

	n	a	b	R <sup>2</sup>	NSE	MAE	RMSE	RSR	MPE
Unit	-	-	-	-	-	mm d <sup>-1</sup>	mm d <sup>-1</sup>	-	%
Case A	1966	0.95	-0.09	0.93	0.93	0.48	1.4	0.27	7.10 (n=1312)
Case B	1836	0.91	0.39	0.87	0.95	0.35	0.57	0.23	7.10 (n=1285)
Case C	1699	0.91	0.38	0.86	-1012	0.34	0.56	31.83	7.10 (n=1284)

For all the three cases, the p-value was 0.05 (not shown). It indicated that existence of statistically significant difference between daily  $E_p$  at 7 a.m. and daily  $E_p$  at midnight values at 5% significance level could not be concluded. In all cases, the MPE was around 7.1%; which means that  $E_p$  at 7 a.m. values were relatively larger by 7.1% as compared to  $E_p$  at midnight values (see Table 3).

For case A, three 'goodness-of-fit' measures ( $R^2$ , NSE, and RSR) showed 'very good' relationship between daily  $E_p$  values at midnight and at 7 a.m. (see Table 3). However, the RMSE was comparatively the biggest. Moreover, in case A (see Table 2) daily  $E_p$  at 7 a.m. and at midnight resulted in 43 and 84 days with  $E_p \leq -0.5 \text{ mm d}^{-1}$  and in 28 and 21 days with  $E_p \geq 7.2 \text{ mm d}^{-1}$ , respectively. Thus, comparatively, using case B or case C was better than using case A. As compared to case C, Case B had the advantage of using more values with better NSE and RSR. Hence, it can be concluded that case B was the best.

Note that in this article the minimum possible  $E_p$  from Class-A pan for the climate condition of Germany is limited to  $\geq -0.5 \text{ mm d}^{-1}$  in an assumption that there could be a maximum condensation of up to  $0.5 \text{ mm d}^{-1}$ . Similarly, the maximum possible  $E_p$ , as calculated using Eq. 6, is limited to  $\leq 7.2 \text{ mm d}^{-1}$  [5-12].

#### Filling missing values of $E_p$ at 7 a.m.

Case B resulted in a regression equation (see Eq. 7 and Figure 4) which is used for filling the missing values of daily  $E_p$  at 7 a.m. by using daily  $E_p$  at midnight values as given below:

$$E_p \text{ at 7 a.m. (in mm d}^{-1}) = 0.908 E_p \text{ at midnight (in mm d}^{-1}) + 0.385 \text{ mm d}^{-1} \quad \dots \quad (7)$$

In Eq. 7 a zero offset was expected; however, an offset of 0.385 had resulted. The cause for non-zero offset might be instrument error; the sensitivity of the pressure sensor instrument for same pressure (depth of pan water) at 7 a.m. and at midnight

might have slight difference. Assuming no instrument error, for example for case B, out of 1836 days,  $E_p$  at 7 a.m. was larger than  $E_p$  at midnight for 1184 days (64.5% of the days). Note also that

$R^2 = 0.87$  in Eq. 7 indicates that daily  $E_p$  at 7 a.m. explains 87 % of the variability in the observed data (daily  $E_p$  at midnight values).

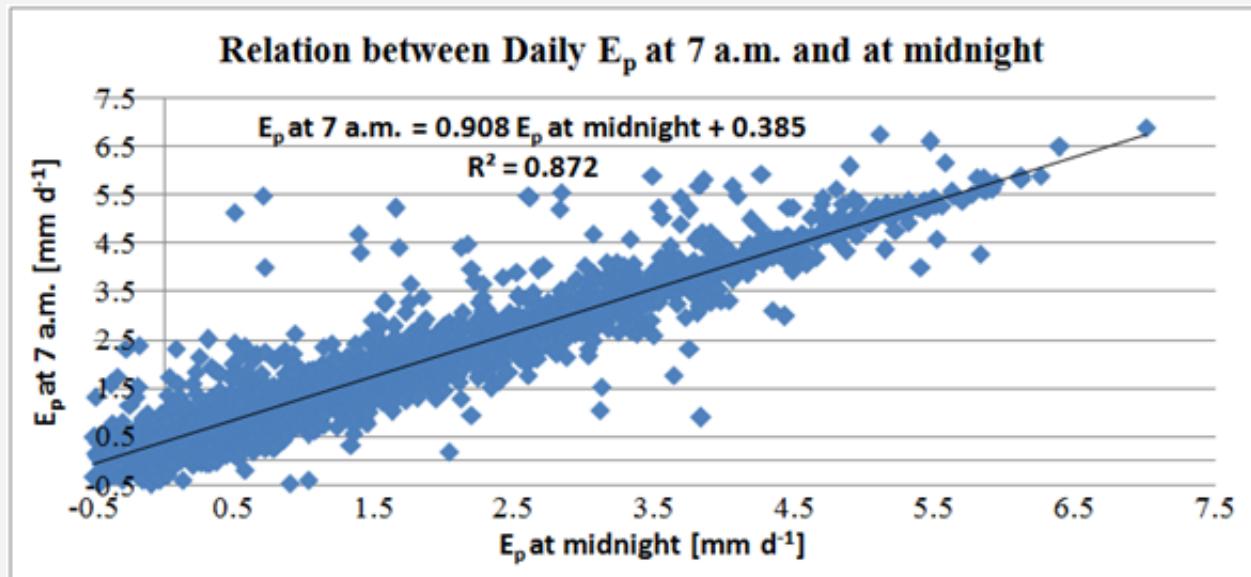


Figure 4: Relation between daily  $E_p$  at 7 a.m. and daily  $E_p$  at midnight at Tharandt for case B ( $-0.5 \leq E_p \leq 7.2$ ) from 2004 to 2013 ( $n = 1836$ ).

Therefore, out of 2145 daily values of  $E_p$  at 7 a.m. the values which are missing and were not in the range between -0.5 and 7.2 mm  $d^{-1}$  are corrected using 1971  $E_p$  at midnight values (Table

2). Accordingly, 24 values of  $E_p$  at 7 a.m. are filled or replaced using Eq. 7 while other 47 values are omitted and 2098 (2145 minus 47) values of  $E_p$  at 7 a.m. are used for next analyses.

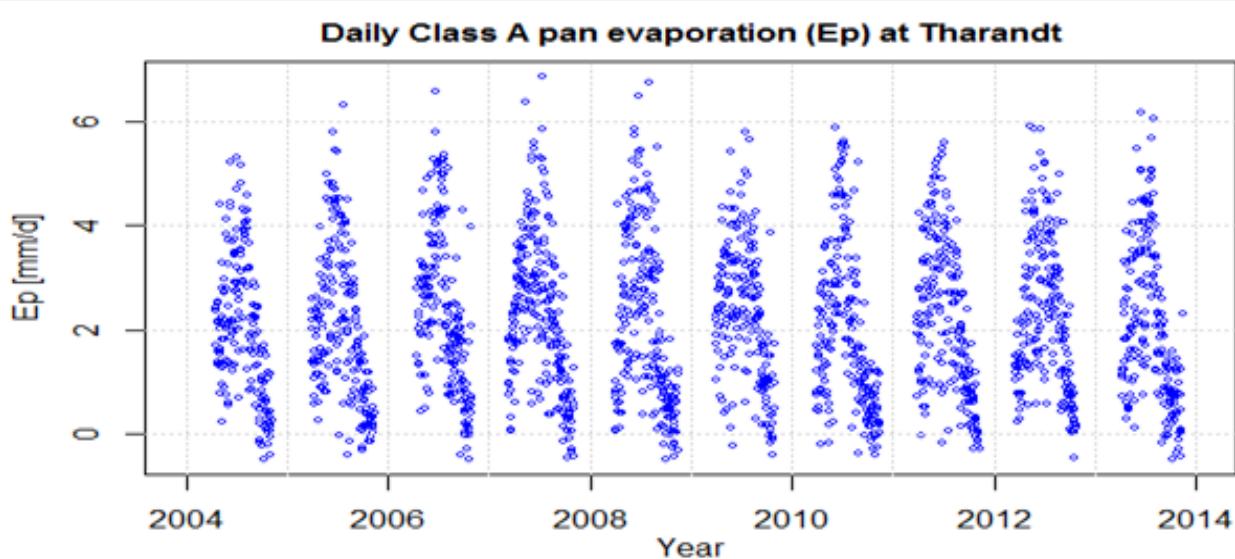


Figure 5: Daily  $E_p$  from March/April to October/November of each year at Tharandt.

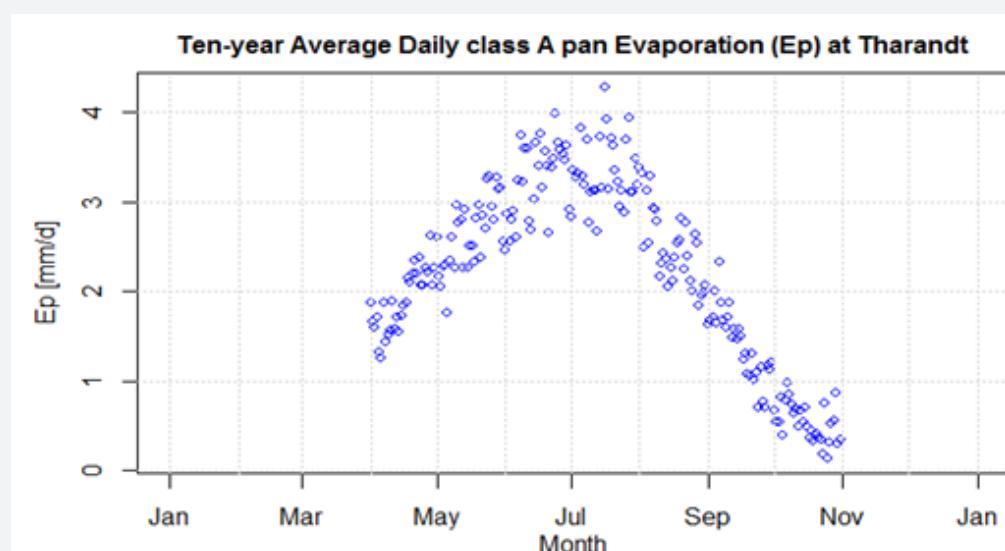
#### Estimated class-A pan evaporation ( $E_p$ )

As discussed in the above sections, daily Class-A pan evaporation ( $E_p$ ) is calculated using daily  $E_p$  at 7 a.m. values for  $n = 2098$  days from March/April to October/November.

Accordingly, the daily  $E_p$  resulted in average, extreme maximum and extreme minimum values of approximately 2.16, 6.87 and -0.50 mm  $d^{-1}$ , respectively. Throughout the 2004 to 2013 period, it was above 6 mm  $d^{-1}$  for only 8 days (see Figure 5).

The ten-year daily average  $E_p$  calculated from April to October (see Figure 6) was approximately between 3 and 4mm  $d^{-1}$  from mid of May to mid of August except for one day in July

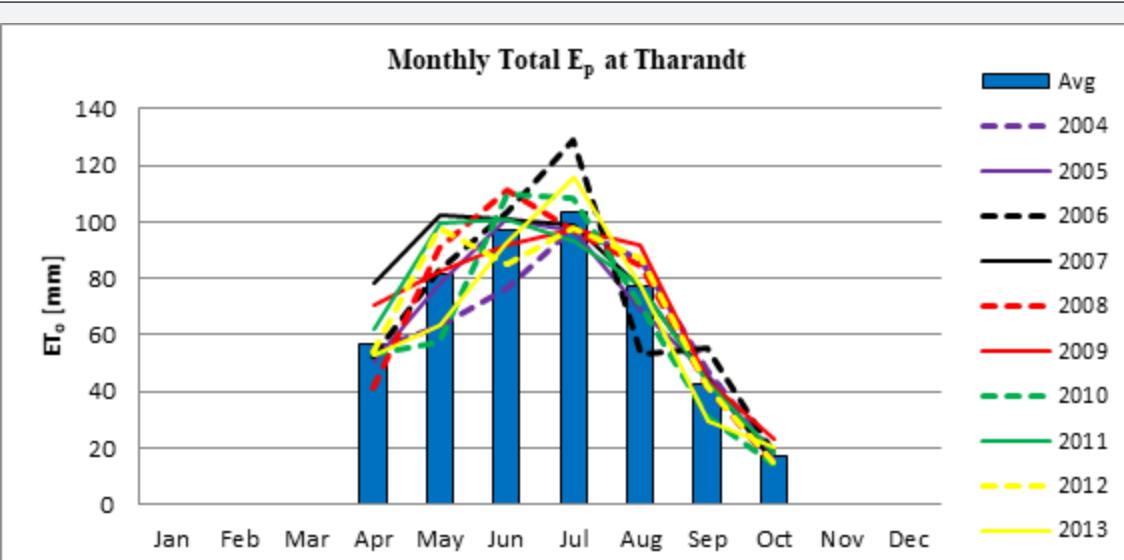
where it was around 4.28mm  $d^{-1}$ . In the rest of the months it was between 3 and 1mm  $d^{-1}$  except from mid of September to October where it declined to between 1 & 0.14mm  $d^{-1}$ .



**Figure 6:** Ten-year (from 2004 to 2013) average of daily  $E_p$  at Tharandt from April to October.

As shown below in Figure 7, the monthly total  $E_p$  calculated from April to October had the highest value in July (103.2mm) followed by June (97.3mm), May (81.8mm) and August (77.4mm). The least value of  $E_p$  was in October (17.3mm)

followed by September (42.4mm) and April (57.0mm). Figure 7 also shows that the peak value was in July for five years, in June for four years and in May for one year.

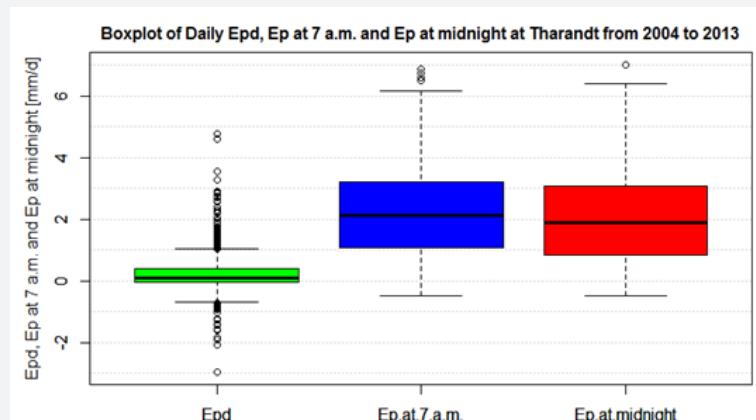


**Figure 7:** Yearly and Average of ten years Monthly Total  $E_p$  at Tharandt from 2004 to 2013.

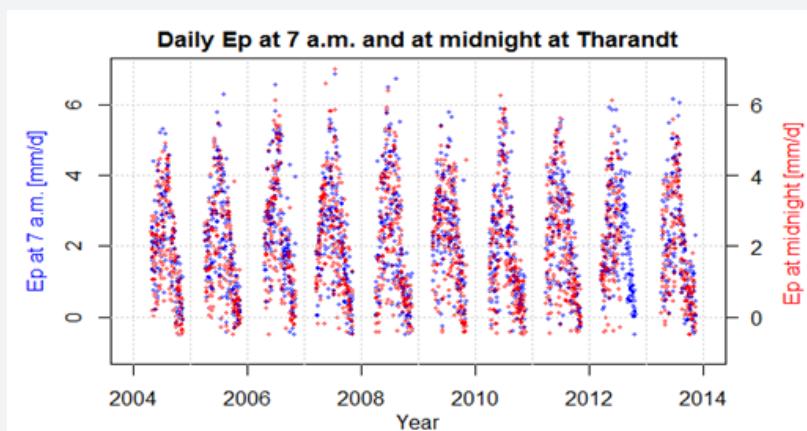
### Understanding $E_{pd}$ (for case B)

$E_{pd}$  (daily  $E_p$  at 7 a.m. minus daily  $E_p$  at midnight) means a next day (Day n+1) 'night time'  $E_p$  minus a previous day (Day n) 'night time'  $E_p$  (see Figure 1). Note that 'night time' is used in this thesis to represent the time from midnight to 7 a.m.

Like Table 3, Figure 8 & 9 graphically show that there was a general 'good' relationship between  $E_p$  at 7 a.m. and  $E_p$  at midnight;  $E_{pd}$  was between  $\pm 1$  mm  $d^{-1}$  for majority of the days. Daily  $E_{pd}$  was  $> 1$  and  $< -1$  mm  $d^{-1}$  for 107 and 14 days, respectively.



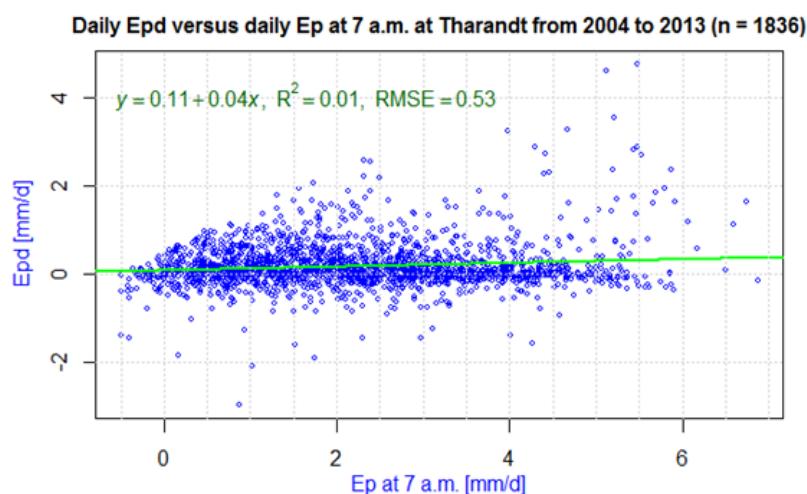
**Figure 8:** Box plot of daily values of  $E_{pd}$ ,  $E_p$  at 7 a.m. and  $E_p$  at midnight at Tharandt from 2004 to 2013 ( $n = 1836$ ).



**Figure 9:** Time series of daily  $E_p$  at 7 a.m. ( $n = 1998$ ) and at midnight ( $n = 1866$ ) at Tharandt from 2004 to 2013.

In Figure 10 the regression line shows only a slight increment in  $E_{pd}$ . The increase was very small; from about 0.1mm to 0.4mm. That means an increase of around 0.3mm per

10 years. Moreover,  $R^2$  was too low. Thus, the trend (existence of a systematic increase or decrease) of  $E_{pd}$  can be neglected.



**Figure 10:** Daily  $E_{pd}$  versus daily  $E_p$  at 7 a.m. at Tharandt from 2004 to 2013 ( $n=1836$ ).

If all other conditions are constant or if instrument and calculation errors are negligible,  $E_{pd}$  shows the 'night time' difference of  $E_p$  (in  $\text{mm d}^{-1}$ ) between two consecutive days. Also, because:

1. there was no a systematic significant trend, shift or lag (Figure 9 and Figure 10);
2. daily  $E_p$  at 7 a.m. and at midnight have good correlation ( $R^2 = 0.87$ ); and
3.  $E_{pd} > 1.5 \text{ mm d}^{-1}$  for 48 days and  $< -1.5 \text{ mm d}^{-1}$  for 6 days; it can be concluded that the 'night time' daily  $E_p$  had to be greater than  $1.0 \text{ mm d}^{-1}$  for 54 days.

## **Conclusion**

The accuracy of the self-recording ten-minute and daily water level measurements from Class-A pan at Tharandt from 2004 to 2013 could be considered as very good. However, the accuracy of the pressure sensor instrument which is used to automatically measure Class-A pan water level shall be carefully checked for the available daily and ten-minute measurements; it might have sensitivity to other than pressure or water depth difference in the pan. Hence, the sensitivity of the pressure sensor instrument at 7 a.m. and at midnight for same pressure (depth of pan water) might have slight difference.

Missing values of daily  $E_p$  at 7 a.m. can be filled using  $0.908 \times \text{daily } E_p \text{ at midnight values} + 0.385$  (Eq. 7). In Eq. 7, the cause for non-zero offset (0.385) could be instrument error. Assuming no instrument error, at Tharandt from 2004 to 2013, out of 1836 days,  $E_{pd}$  was larger than zero for 1184 days (64.5% of the days). It is also concluded that the 'night time' daily  $E_p$  had to be  $\geq 1.0 \text{ mm d}^{-1}$  for 54 days. The existence of 'night time'  $E_p$  might have made the comparison of  $E_p$  at 7 a.m. and  $E_p$  at midnight a bit complicated.

The sensitivity of the automatic Class-A pan water level measuring instrument to other than pressure (water depth difference in the pan) must be checked. If there was no measurement error and if 'night time'  $E_p$  is negligible, then the question: 'Why  $E_p$  at 7 a.m. is greater than  $E_p$  at midnight for majority (65%) of the days?' may require further study.

## **Acknowledgement**

For every good thing, I praise GOD and GOD'S MOTHER above all. I particularly thank Virgin Mary or 'Tsadkane Mariam' (ጥሃቅና ማርያም) monastery of Ethiopia for I get healings!

Very special thanks to my official supervisors Prof. Dr. Christian Bernhofer and Dr. Uta Moderow for their excellent supervision, for providing me all the needed data in advance,

and for their friendly approach throughout my research. Very special thanks to Mr. Endalkachew Bekele for giving me valuable guidance and for editing the text for publishing. I also thank Mr. Thomas Pluntke, Mr. Philipp Körner and their colleague(s) for their support and for providing me meteorological data.

I also want to thank all my elementary and high school teachers especially my mathematics teachers Mrs. Abebech ('Tiye Abebech') and Mr. Abebe ('Gash Abe'). I also want to use this chance to thank all kind persons who (have) supported me in my life.

Last but the best, I would like to thank my wife and my family and friends for their crucial support and for sharing love and happiness throughout my life.

## **References**

1. Dingman SL (2002) Physical Hydrology. (2<sup>nd</sup> edn), Upper Saddle River, NJ., Prentice Hall.
2. McMahon TA, Peel MC, Lowe L, Srikanthan R, McVicar TR (2013) Estimating actual, potential, reference crop and pan evaporation using standard meteorological data: a pragmatic synthesis. Hydrol Earth Syst Sci 17: 1331-1363.
3. Dingman SL (1994) Physical Hydrology. (3<sup>rd</sup> edn).
4. Allen RG, Pereira LS, Raes D, Smith M (1998) Crop evapotranspiration: Guidelines for computing crop water requirements. FAO Irrig. and Drain. Paper No. 56, Food and Agricultural Organization of the United Nations, Rome, Italy.
5. ASCE-EWRI (2005) The ASCE Standardized Reference Evapotranspiration Equation. Task Committee (TC) on Standardization of Reference Evapotranspiration. In: Allen RG, Walter IA, Elliott R, Howell T, Itenfisu D, et al. (Eds.), Environmental and Water Resources Institute (EWRI) of the American Society of Civil Engineers (ASCE).
6. Filho DBF, Paranhos R, Rocha EC, Batista M, Silva JA, et al. (2013) When is statistical significance not significant? *Brazilian political science review*.
7. Irmak S, Haman DZ, Jones JW (2002) Evaluation of Class-A Pan Coefficients for Estimating Reference Evapotranspiration in Humid Location. J Irrig Drain Eng ASCE 128(3): 153-159.
8. Khan AU, Hildreth WB (2003) Case studies in public budgeting and financial management. New York, N.Y: Marcel Dekker, USA.
9. Legates DR, McCabe GJ (1999) Evaluating the use of "goodness-of-fit" measures in hydrologic and hydroclimatic model validation. Water Resources Research 35(1): 233-241.
10. Lhomme JP (2016) Towards a rational definition of potential evaporation. *Hydrology and Earth System Sciences* 1(2): 257-264.
11. Mohammadi M, Ghahraman B, Davary K, Liaghat AM, Bannaya M (2012) Pan coefficient ( $K_p$ ) estimation under uncertainty on fetch. Meteorol Atmos Phys 117(1-2): 73-83.
12. Moriasi DN, Arnold GJ, Van Liew MW, Bingner RL, Harmel RD, et al. (2007) Model Evaluation Guidelines for Systematic Quantification of Accuracy in Watershed Simulation. American Society of Agricultural and Biological Engineers 50(3): 885-900.



This work is licensed under Creative  
Commons Attribution 4.0 License  
DOI: [10.19080/IJESNR.2019.19.556003](https://doi.org/10.19080/IJESNR.2019.19.556003)

**Your next submission with Juniper Publishers  
will reach you the below assets**

- Quality Editorial service
- Swift Peer Review
- Reprints availability
- E-prints Service
- Manuscript Podcast for convenient understanding
- Global attainment for your research
- Manuscript accessibility in different formats  
**( Pdf, E-pub, Full Text, Audio)**
- Unceasing customer service

Track the below URL for one-step submission

<https://juniperpublishers.com/online-submission.php>